

INTRODUCTION

LaMEM (Lithosphere and Mantle Evolution Model) is a 3D, open-source, thermo-mechanical numerical code to simulate crustal and lithospheric deformation. The code is written on top of PETSc and is based on a marker-and-cell approach combined with a staggered finite difference (FDSTAG) discretization in space, which is a stable and very efficient technique to solve the (nearly) incompressible Stokes equations that does not suffer from spurious pressure modes or artificial compressibility (a typical feature of low-order finite element techniques) and has the same accuracy as FE methods for typical geodynamic applications. Here we discuss some issues related to the accuracy of the method, convergence of non-linear solvers and show a few recent modelling applications.

Non-linear visco-elasto-plastic rheology

Effective creep viscosity plot

Stress update
 $\tau_{ij} = 2\eta \dot{\epsilon}_{ij}$
 $\dot{\epsilon}_{ij} = \frac{1}{2}(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i})$
 $\eta = \frac{A}{B} \exp\left[-\frac{E_a}{RT}\right] \exp\left[-\frac{p}{p_0}\right] \exp\left[\frac{v}{v_0}\right]$

Effective strain rate and invariant
 $\dot{\epsilon}_{ij} = \frac{1}{2}(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i})$
 $\dot{\epsilon} = \sqrt{\frac{2}{3} \dot{\epsilon}_{ij} \dot{\epsilon}_{ij}}$

Stress rotation terms
 $\tau_{ij} = \tau_{ij}^n + \Delta t (w_{ik} \tau_{kj}^n - \tau_{ik}^n w_{kj})$
 $w_{ij} = \frac{1}{2}(\frac{\partial v_i}{\partial x_j} - \frac{\partial v_j}{\partial x_i})$

Drucker-Prager Effective viscosity
 $\tau_{ij} = \mu P + c \cdot \sigma^{-1} \min\left(\frac{\partial v_i}{\partial x_j} + \frac{\partial v_j}{\partial x_i}, \frac{\partial v_i}{\partial x_i} + \frac{\partial v_j}{\partial x_j}\right)$

Diffusion, dislocation and Peierls constants
 $A_i = B_i \exp\left[-\frac{E_{a,i}}{RT}\right]$
 $A_n = B_n \exp\left[-\frac{E_{a,n}}{RT}\right]$
 $A_p = \frac{B_p}{(v_p)^2} \exp\left[-\frac{E_{a,p}}{RT}\right] (1-\gamma)^2$

Peierls effective exponent
 $\beta = \frac{E_{a,p}}{RT} (1-\gamma)^{-1} \gamma^2$

Finite element Jacobian derivation (matrix notation)

Stress and strain rate vectors
 $\tau = [\tau_{11}, \tau_{22}, \tau_{33}, \tau_{12}, \tau_{13}, \tau_{23}]^T$
 $\dot{\epsilon} = [\dot{\epsilon}_{11}, \dot{\epsilon}_{22}, \dot{\epsilon}_{33}, \dot{\epsilon}_{12}, \dot{\epsilon}_{13}, \dot{\epsilon}_{23}]^T$

Nonlinear residuals
 $r^T = \int_{\Omega} B^T (r - mP) - N_{ij}^T p_{ij} d\Omega$

Jacobian blocks
 $J^T = \int_{\Omega} B^T Q T Q^T B d\Omega$
 $J^P = - \int_{\Omega} B^T (m - pQ) N_P d\Omega$
 $J^W = - \int_{\Omega} N_{ij}^T m^T B d\Omega$

Visco-elastic nonlinear parameter
 $\beta = \frac{1}{2} \left(\frac{1}{n} - 1\right) \frac{v}{v_0} + \frac{1}{2} \left(\frac{1}{n} - 1\right) \frac{v}{v_0}$

Plastic nonlinear parameter
 $\beta = -\frac{1}{2}$

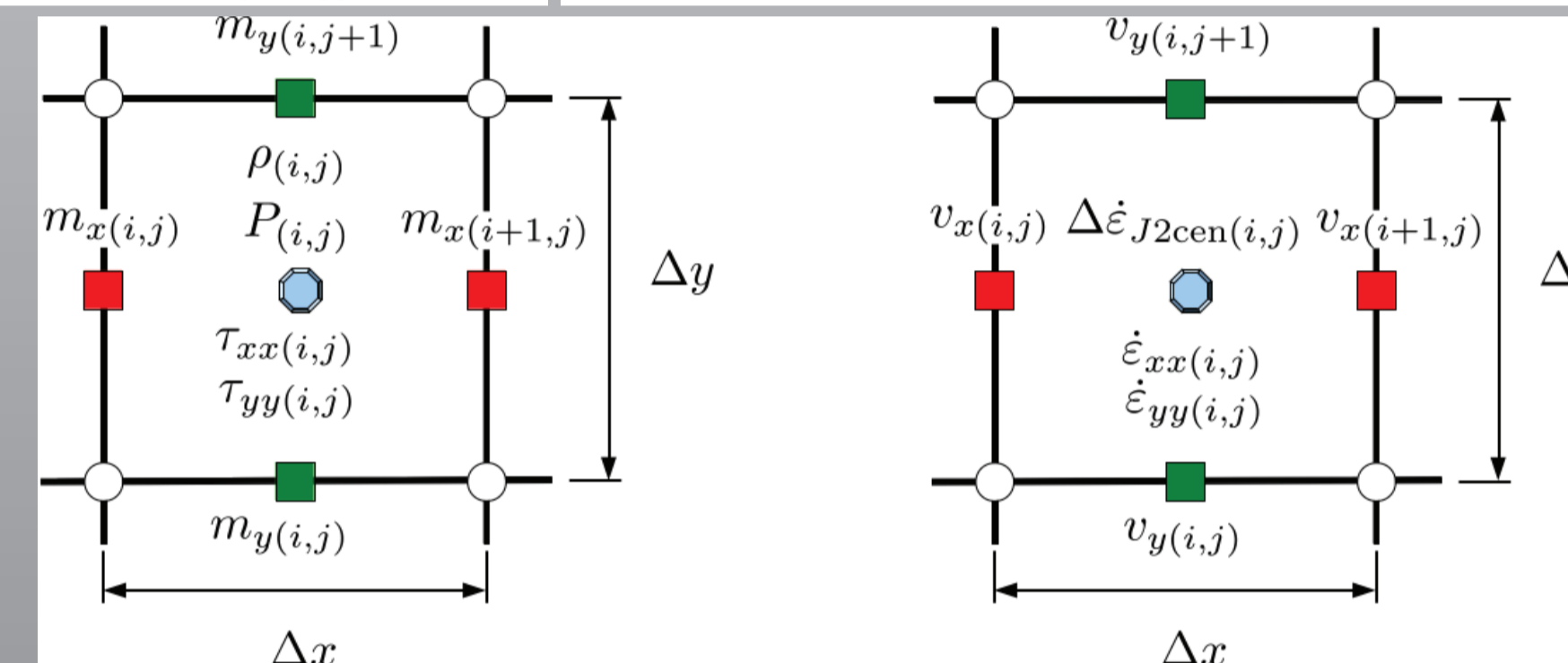
Normalized flow vector
 $g = \frac{\dot{\epsilon}}{\dot{\epsilon}}$

Pressure projection vector
 $m = [1 \ 1 \ 1 \ 0 \ 0 \ 0]^T$

Projection matrix and projection
 $Q = \begin{bmatrix} 2 & 0 & 0 & 0 & 0 & 0 \\ 0 & 2 & 0 & 0 & 0 & 0 \\ 0 & 0 & 2 & 0 & 0 & 0 \\ 0 & 0 & 0 & 1 & 1 & 0 \\ 0 & 0 & 0 & 1 & 1 & 0 \\ 0 & 0 & 0 & 0 & 0 & 1 \end{bmatrix}$
 $r = Qr, \quad \epsilon = Q^T \epsilon$

Jacobian iteration
 $\begin{bmatrix} \Delta P_{i+1} \\ \Delta \tau_{i+1} \end{bmatrix} = - \begin{bmatrix} J^P & J^W \\ J^T & J^W \end{bmatrix}^{-1} \begin{bmatrix} r_i \\ \tau_i \end{bmatrix}$

Tangent matrix
 $T = 2\eta (I + \beta g g^T)$



Staggered finite difference Jacobian derivation (single component example)

Momentum residual contribution
 $\Delta m_{x(i,j)} = -\frac{\partial \tau_{xx(i,j)}}{\partial x} + \frac{\partial \tau_{xy(i,j)}}{\partial y} + \rho g_x$
 $\Delta m_{y(i,j)} = -\frac{\partial \tau_{xy(i,j)}}{\partial x} + \frac{\partial \tau_{yy(i,j)}}{\partial y} + \rho g_y$

Strain rate derivatives
 $\frac{\partial \dot{\epsilon}_{xx(i,j)}}{\partial x} = \frac{\partial^2 u_{x(i,j)}}{\partial x^2} - \frac{\partial^2 v_{x(i,j)}}{\partial x \partial y}$
 $\frac{\partial \dot{\epsilon}_{yy(i,j)}}{\partial y} = \frac{\partial^2 v_{y(i,j)}}{\partial y^2} - \frac{\partial^2 u_{y(i,j)}}{\partial x \partial y}$

Second invariant contribution derivatives
 $\frac{\partial \dot{\epsilon}}{\partial x} = \frac{\partial^2 u_{x(i,j)}}{\partial x^2} + \frac{\partial^2 v_{y(i,j)}}{\partial y^2}$
 $\frac{\partial \dot{\epsilon}}{\partial y} = -\frac{\partial^2 u_{x(i,j)}}{\partial x \partial y} + \frac{\partial^2 v_{y(i,j)}}{\partial x \partial y}$

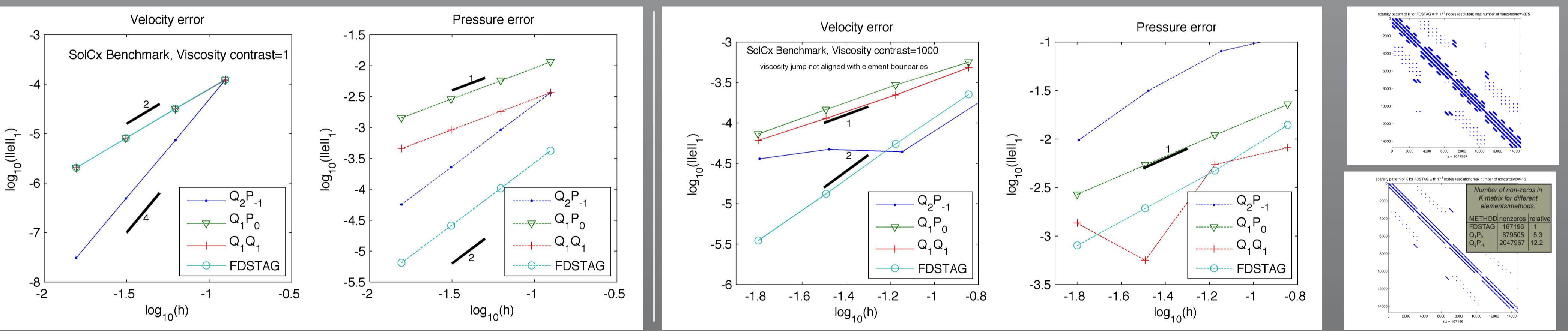
Residual derivatives (velocity)
 $\frac{\partial \Delta m_{x(i,j)}}{\partial u_{x(i,j)}} = -\frac{\partial^2}{\partial x^2}$
 $\frac{\partial \Delta m_{x(i,j)}}{\partial v_{x(i,j)}} = \frac{\partial^2}{\partial x \partial y}$

Residual derivatives (pressure)
 $\frac{\partial \Delta m_{x(i,j)}}{\partial p} = \frac{\partial}{\partial x}$
 $\frac{\partial \Delta m_{y(i,j)}}{\partial p} = \frac{\partial}{\partial y}$

Contributions from neighbor points are assembled in the same way

The full, discretized, Jacobian increases the stencilwidth for FDSTAG. That is why we implement the action of the matrix-vector product in a matrix-free manner.

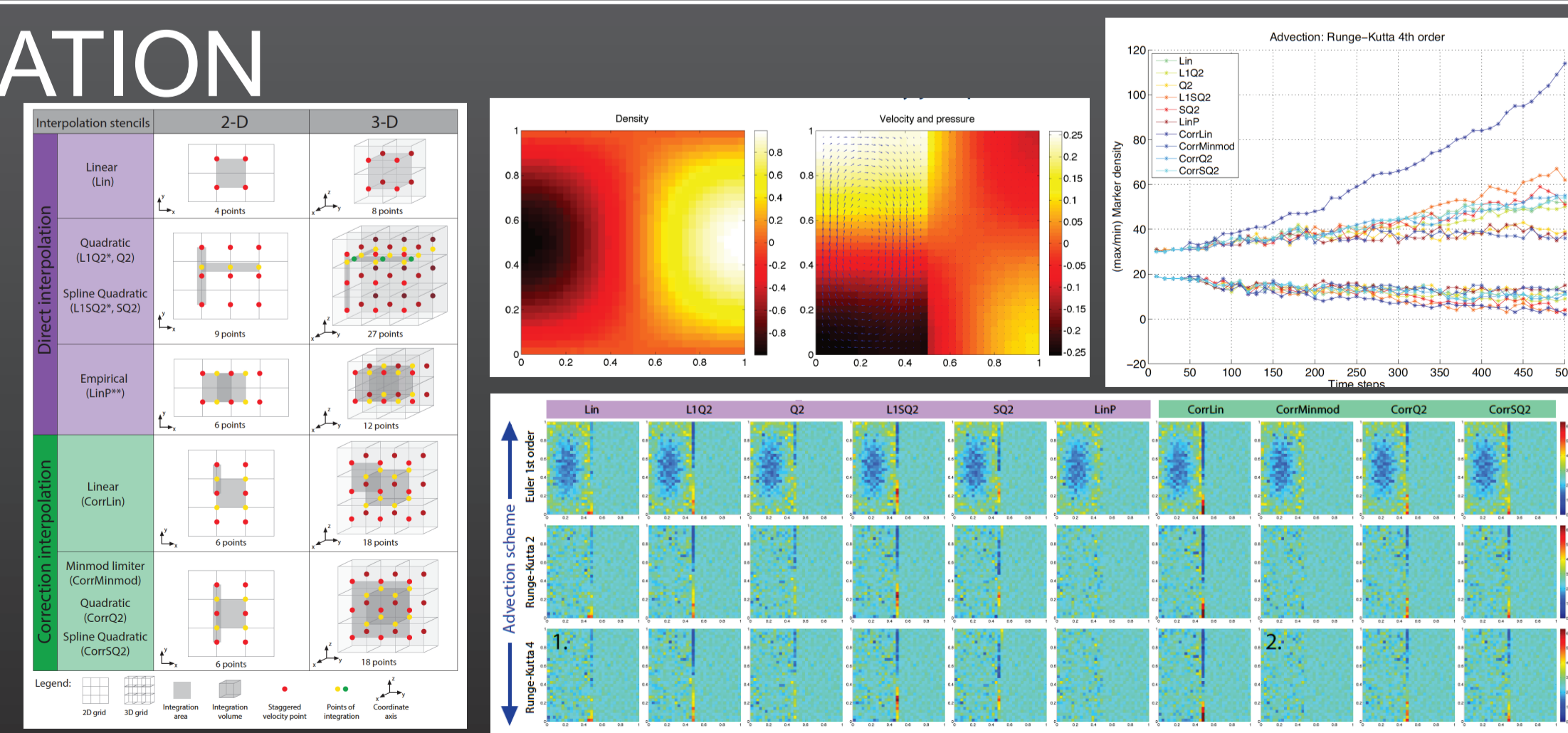
ACCURACY & MEMORY COSTS: FEM vs. FDSTAG



Some claim that higher-order finite element codes are more accurate than staggered finite difference discretizations. As the results show, this is only true if the jumps in viscosity are *exactly* aligned with the boundaries of the finite element. In typical evolving geodynamic setups (e.g., subduction), jumps do occur within an element either due to the marker distribution or due to non-linearities (e.g., plastic localization). Under these conditions, all methods are first order accurate [Deubelbeiss & Kaus, 2008; Thielmann et al. 2014]. Yet, low order finite elements (Q1P0) are unstable (wrong pressure), whereas the staggered finite difference is a stable discretization. The stabilized Q1Q1 element is compressible which is a problem in setups with a free surface. The Q2P1 element is stable but expensive (in 3D) as it requires up to 12 times more memory for the same number of nodes, which makes every matrix-vector product an order or magnitude slower. The staggered finite difference method is thus a cheap and stable low-order discretization method for typical geodynamic problems. A potential disadvantage is that the jacobian is more complicated to implement and increases the stencil width of the discretization. Yet, our results show it is typically not necessary to assemble the jacobian but that, instead, a matrix-free implementation is sufficient in many cases. Benchmarks show that a free surface can be successfully represented by a sticky air layer [Cramer et al., GJI 2012], particularly if the 'air' is implemented by a stress-free internal boundary condition [Duret et al., GJI 2016].

CONSERVATIVE MAC INTERPOLATION

The marker-and-cell approach involves frequent interpolations of velocities from nodal points to markers. This introduces interpolation errors as the velocity at markers is no longer guaranteed to be divergence-free which becomes noticeable with time, as areas devoid of particles start appearing (which require particle injection/removal). A number of workarounds have been proposed recently, yet it is unclear which of those works best for a geodynamics application in particular combined with FDSTAG. We have adapted them for FDSTAG formulations and found the MINMOD method, combined with a 4th order Runge Kutta time integration scheme to be superior [Püsök et al., submitted].



Coupled vs. decoupled MG

Coupled vs. decoupled MG

$P_c = \begin{pmatrix} K & G \\ D & -\frac{1}{\eta} I \end{pmatrix}$ (COUPLED) Galerkin MG applied to full matrix

$P_u = \begin{pmatrix} K & G \\ & -\frac{1}{\eta} I \end{pmatrix}$ (UNCOUPLED) Galerkin MG applied to K block only

Restriction
 $r_{coarse} = R r_{fine}$

Prolongation
 $x_{fine} = P x_{coarse}$

Coarsening
 $A_{coarse} = R A_{fine} P$

GALERKIN COARSENING

64³ nodes

Viscosity contrast	Coupled multigrid # outer KSP it	Block MG # outer KSP it	Time [s]	Total solve [s]
1	7	6	6.1	3.9
10	10	8	11	6
100	15	11	17	10.2
1000	36	27	45	23
10000	114	82	154	80.5
1.00E+05	378	266	585	297

Subduction of lithospheric plates with viscoplastic rheology

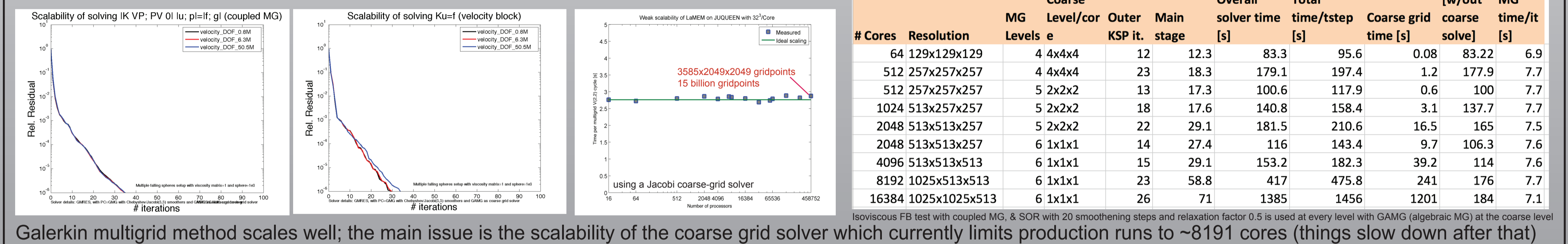
Resolution	# SNES (nonlinear)	# KSP it (total)	Time per timestep [s]
128x32x32	2	40	12
	2	100	24

4 GMG levels

Resolution	# SNES (nonlinear)	# KSP it	Time [s]
256x64x64	2	30	145
	5	250	759

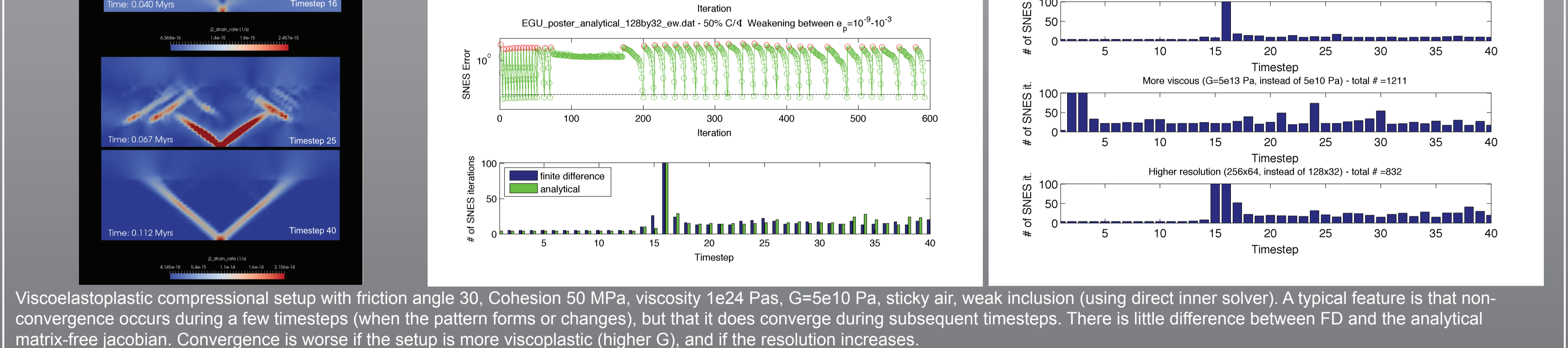
One can do Galerkin multigrid only for the velocity block, or for both velocity and pressure. Which one is best? For simple, linear viscous falling sphere setups there is little difference. For realistic viscoplastic setups, coupled wins.

PARALLEL SCALABILITY



Galerkin multigrid method scales well; the main issue is the scalability of the coarse grid solver which currently limits production runs to ~8191 cores (things slow down after that)

PLASTICITY



Viscoelastoplastic compressional setup with friction angle 30, Cohesion 50 MPa, viscosity 1e24 Pas, sticky air, weak inclusion, weak inclusion. A typical feature is that non-convergence occurs during a few timesteps (when the pattern forms or changes), but that it does converge during subsequent timesteps. There is little difference between FD and the analytical matrix-free jacobian. Convergence is worse if the setup is more viscoplastic (higher G), and if the resolution increases.

ONGOING WORK

- We are currently focussing on a few topics
- (1) Better integrating the solvers with the PETSc framework, by using a DMSHELL for the FDSTAG
 - (2) Coupling of the vep Stokes code with two-phase melt migration.
 - (3) Adding an explicit poro-visco-elasto-plastic solver to simulate hydrofracking
 - (4) Inflow/outflow boundary conditions.

CONCLUSIONS

- LaMEM is a parallel 3D marker-and-cell based deformation code suitable to model lithospheric processes.
- In typical geodynamic cases, FDSTAG is relative powerful (and cheap) alternative to finite-element calculations.
- The code can perform production runs on massively parallel machines.
- Conservative marker-in-cell interpolation was added & improves marker distributions.
- Coupled multigrid outperforms decoupled multigrid for nonlinear setups.
- Elasticity helps convergence for setups with brittle plastic rheology, even though more work on this topic is required to (either different stabilizing rheology or better convergent nonlinear algorithms).

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